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Evaluation of operational on-line-coupled regional air quality models over Europe and North America in the context of AQMEII phase 2. Part I: Ozone

Ulas Im, Roberto Bianconi, Efisio Solazzo, Ioannis Kioutsioukis, Alba Badia, Alessandra Balzarini, Rocío Baró, Roberto Bellasio, Dominik Brunner, Charles Chemel, Gabriele Curci, Johannes Flemming, Renate Forkel, Lea Giordano, Pedro Jiménez-Guerrero, Marcus Hirtl, Alma Hodzic, Luka Honzak, Oriol Jorba, Christoph Knote, Jeroen J.P. Kuenen, Paul A. Makar, Astrid Manders-Groot, Lucy Neal, Juan L. Pérez, Guido Pirovano, George Pouliot, Roberto San Jose, Nicholas Savage, Wolfram Schroder, Ranjeet S. Sokhi, Dimiter Syrakov, Alfreida Torian, Paolo Tuccella, Johannes Werhahn, Ralf Wolke, Khairunnisa Yahya, Rahela Zabkar, Yang Zhang, Junhua Zhang, Christian Hogrefe, Stefano Galmarini



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- 6 Balzarini^d, Rocío Baró^e, Roberto Bellasio^b, Dominik Brunner^f, Charles Chemel^g, Gabriele
- 7 Curci^h, Johannes Flemmingⁱ, Renate Forkel^j, Lea Giordano^f, Pedro Jiménez-Guerrero^e,
- 8 Marcus Hirtl^k, Alma Hodzic¹, Luka Honzak^m, Oriol Jorba^c, Christoph Knote¹, Jeroen J.P.
- 9 Kuenenⁿ, Paul A. Makar^o, Astrid Manders-Grootⁿ, Lucy Neal^p, Juan L. Pérez^q, Guido
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- 12 Ralf Wolke^s, Khairunnisa Yahya^u, Rahela Zabkar^{m,v}, Yang Zhang^u, Junhua Zhang^o, Christian
- 13 Hogrefe^r, Stefano Galmarini^{a*}

14

15	a.	European Commission, Joint Research Centre, Institute for Environment and
16		Sustainability, Air and Climate Unit, Ispra (Italy).
17	b.	Enviroware srl, Concorezzo (MB), Italy.
18	с.	Earth Sciences Department, Barcelona Supercomputing Center (BSC-CNS),
19		Barcelona, Spain.
20	d.	Ricerca sul Sistema Energetico (RSE SpA), Milano, Italy
21	e.	University of Murcia, Department of Physics, Physics of the Earth. Campus de
22		Espinardo, Ed. CIOyN, 30100 Murcia, Spain.
23	f.	Laboratory for Air Pollution and Environmental Technology, Empa, Dubendorf,
24		Switzerland.
25	g.	Centre for Atmospheric & Instrumentation Research, University of Hertfordshire,
26		College Lane, Hatfield, AL10 9AB, UK.
27	h.	Department of Physical and Chemical Sciences, Center of Excellence for the forecast
28		of Severe Weather (CETEMPS), University of L'Aquila, L'Aquila, Italy.
29	i.	ECMWF, Shinfield Park, RG2 9AX Reading, United Kingdom.
30	j.	Karlsruher Institut für Technologie (KIT), Institut für Meteorologie und
31		Klimaforschung, Atmosphärische Umweltforschung (IMK-IFU), Kreuzeckbahnstr. 19,
32		82467 Garmisch-Partenkirchen, Germany.
33	k.	Section Environmental Meteorology, Division Customer Service, ZAMG -
34		Zentralanstalt für Meteorologie und Geodynamik, 1190 Wien, Austria.
35	1.	National Center for Atmospheric Research, Boulder, CO, US.
36	m.	Center of Excellence SPACE-SI, Ljubljana, Slovenia.
37	n.	Netherlands Organization for Applied Scientific Research (TNO), Utrecht, The
38		Netherlands.
39	0.	Air Quality Research Section, Atmospheric Science and Technology Directorate,
40		Environment Canada, 4905 Dufferin Street, Toronto, Ontario, Canada.
41	p.	Met Office, FitzRoy Road, Exeter, EX1 3PB, United Kingdom.
42	q.	Environmental Software and Modelling Group, Computer Science School - Technical
43		University of Madrid, Campus de Montegancedo - Boadilla del Monte-28660, Madrid,
44		Spain.

45 46	r.	Emissions and Model Evaluation Branch, Atmospheric Modeling and Analysis Division/NERL/ORD, Research Triangle Park, North Carolina, USA.
47	s.	Leibniz Institute for Tropospheric Research, Permoserstr. 15, D-04318 Leipzig,
48	4	Germany.
49 50	t.	National Institute of Meteorology and Hydrology, Bulgarian Academy of Sciences, 66
50	11	Department of Marine, Earth and Atmospheric Sciences, North Carolina State
52	u.	University Raleigh USA
52	v	University, Kaleigh, OSA.
54	۰.	University of Ejubijana, I acuity of Mathematics and Thysics, Ejubijana, Slovena.
55	* Corre	esponding author: S. Galmarini (Stefano.galmarini@irc.ec.europa.eu)
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57	Highlig	ghts
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59	•	Sixteen modeling groups from EU and NA simulated O ₃ for 2010 under AOMEII
60		phase 2
61	•	A general model underestimation of surface O_3 over both continents up to 22%
62	•	Models tend to over/under estimate surface O_3 in all regions during autumn/winter
63	•	Boundary conditions influence O_3 predictions especially during winter and autumn
64	•	Models tend to under-predict high O_3 values that are of concern for policy
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67	Keywo	rds: AQMEII, on-line coupled models, performance analysis, ozone, Europe, North
68	Americ	ca
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70	ABSTI	RACT
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72	The sec	cond phase of the Air Quality Model Evaluation International Initiative (AQMEII)
73	brough	t together sixteen modeling groups from Europe and North America, running eight
74	operati	onal online-coupled air quality models over Europe and North America on common
75	emissio	ons and boundary conditions. With the advent of online-coupled models providing new
76	capabil	ity to quantify the effects of feedback processes, the main aim of this study is to
77	compa	The response of coupled air quality models to simulate levels of O_3 over the two with regions. The simulated enough second continental and sub regional example.
78	contine	ental regions. The simulated annual, seasonal, continental and sub-regional ozone
79 00	larga o	becryptional database from different measurement networks operating in Europe and
00 81	North	America Results show a general model underestimation of the annual surface ozone
82	levels	over both continents reaching up to 18% over Europe and 22% over North America
83	The ob	served temporal variations are successfully reproduced with correlation coefficients
84	larger t	han 0.8 Results clearly show that the simulated levels highly depend on the
85	meteor	ological and chemical configurations used in the models, even within the same
86	modeli	ng system. The seasonal and sub-regional analyses show the models' tendency to
87	overest	timate surface ozone in all regions during autumn and underestimate in winter.
88	Bounda	ary conditions strongly influence ozone predictions especially during winter and
89	autumr	, whereas during summer local production dominates over regional transport. Daily
90	maxim	um 8-hour averaged surface ozone levels below 50-60 μ g m ⁻³ are overestimated by all
91	models	s over both continents while levels over 120-140 μ g m ⁻³ are underestimated, suggesting
92	that mo	odels have a tendency to severely under-predict high O ₃ values that are of concern for
93	air qua	lity forecast and control policy applications.

95

96 1. Introduction

97 Tropospheric ozone (O_3) is an important secondary air pollutant produced by photochemical oxidation of volatile organic compounds (VOC) and carbon monoxide (CO) in the presence of 98 nitrogen oxides (NOx). It has implications on climate and health and therefore its levels are 99 100 subject to regulatory monitoring in Europe (EU) and North America (NA). The regulatory O₃ levels are still exceeded in a number of cities and are especially a concern in growing urban 101 areas (European Environmental Agency, 2013). Air quality models (AQMs) are valuable tools 102 103 to investigate the complex and dynamic interactions between meteorology and chemistry leading to O₃ pollution episodes at multiple temporal and spatial scales. In the last decade, 104 AQM development started shifting from off-line-coupled models where the meteorological 105 forcing for chemistry was produced off-line by a separate meteorological model, to fully-106 coupled online models, which are able to simulate the feedbacks between chemistry and 107 meteorology, taking the advantage of increased computational power (Zhang, 2008; Baklanov 108 et al., 2014). The use of on-line models for O₃ predictions is beneficial, as O₃ not only 109 depends on emissions and chemistry but also on regional transport, clouds, photolysis and 110 vertical mixing in the boundary layer, all of which can be more realistically represented in an 111 on-line model (Wong et al, 2012; Zhang et al, 2013). The impact of aerosols on the radiation, 112 113 and therefore temperatures and photolysis rates, can significantly impact the gas-phase chemistry affecting O₃ and secondary aerosol formation (Kim et al., 2009, 2011). Thus, 114 simulating these feedbacks can lead to more realistic O_3 , NOx and aerosol levels that are 115 relevant to policy applications. The wide use of regional AQMs for supporting policy, 116 abatement strategies and forecasting justifies the increased need for online models, which can 117 simulate feedback mechanisms, and especially account for the effect of aerosols on radiative 118

119 balance and photolysis (e.g. Hodzic et al., 2007).

The Air Quality Model Evaluation International Initiative (AQMEII) served to promote 120 policy-relevant research on regional air quality model evaluation across the atmospheric 121 modeling communities in Europe and North America through the exchange of information on 122 123 current practices and the identification of research priorities (Galmarini and Rao, 2011). As part of this collaboration, standardized observations and model outputs were made available 124 through the ENSEMBLE system (http://ensemble2.jrc.ec.europa.eu/public/) that is hosted at 125 the Joint Research Centre (JRC). This web-interface allows temporal and spatial analyses of 126 individual models as well as their ensemble operators (Bianconi et al., 2004; Galmarini et al., 127 2012). The first phase of AOMEII was focused on the evaluation of off-line coupled 128 atmospheric modelling systems against large sets of monitoring observations over Europe and 129 North America for the year 2006 (Solazzo et al., 2012a,b; Vautard et al., 2012; Solazzo et al., 130 2013; Hogrefe et al., 2014). As summarized in Schere et al. (2012), the intercomparison 131 132 model results for O₃ suggested a strong influence of chemical boundary conditions for ozone, whose bias extends far into the interior of the modelling domains, especially during winter 133 months. The observed variance as well as the daily ozone cycle was underestimated by the 134 majority of models. Night-time, overcast, and stable conditions led to poor model skill in 135 reproducing ozone mixing ratios over both continents. Stable atmospheric boundary layers 136 have been notoriously difficult to simulate in numerical weather prediction models (Holtslag 137

- et al. 2013), but they are highly relevant in the context of air quality modeling. Due to the
- high sensitivity of air pollutants to the representation of stable boundary layers, online
- 140 coupled modeling could be of great use to tackle this problem in the future.
- 141 The second phase of AQMEII extends this model assessment to on-line-coupled air quality
- models. In this study, we analyze O₃ concentrations provided by eight on-line-coupled
- models, which have been run by sixteen independent groups from Europe and North America
- 144 (while a companion study is devoted to the analyses of particulate matter, Im et al., 2014).
- 145 The models made use of the same input emissions and chemical boundary conditions, in an 146 effort to reduce the impact of uncertainties originating from these inputs to model results
- effort to reduce the impact of uncertainties originating from these inputs to model resultsamong different groups. The goal of the study is to evaluate the performances of widely used
- 148 operational on-line coupled models in Europe and North America in simulating O₃ levels on a
- 149 sub-regional and seasonal basis employing an experimental set-up with common
- anthropogenic emission and boundary conditions. The surface levels and vertical profiles
- simulated by the individual models as well as their ensemble mean and median are compared
- 152 with the observational data provided by the ENSEMBLE system.
- 153
- 154 2. Materials and Methods
- 155 2.1. Participating models
- 156 In the context of AQMEII2, twelve modeling groups from EU and four modeling groups from
- 157 NA (Table 1) have applied their modeling systems to simulate hourly O₃ concentrations for
- the year 2010 over the EU and NA continental scale domains (Fig.1). Among all participants,
- seven groups from EU and two groups from NA applied the same model system (WRF-
- 160 CHEM), but with different settings such as different shortwave radiation schemes, gas-phase
- 161 chemical mechanisms and aerosol modules. The WRF-CHEM community applied a common
- horizontal grid spacing of 23 km over Europe and 36 km over North America. Other
- modeling groups applied different grid spacings, ranging from 12×12 km² to $\sim 50 \times 25$ km² as
- seen in Table 1. The simulations were conducted for continental-scale domains of Europe andNorth America covering continental U.S., southern Canada and northern Mexico (Fig.1). To
- 165 North America covering continental U.S., southern Canada and northern Mexico (Fig.1). To 166 facilitate the cross-comparison between models, the participating groups interpolated their
- 167 model output to a common grid with 0.25° resolution for both continents. Model values at
- 168 observation locations were extracted from the original model output files for comparison to
- 169 observations (described below).
- 170 2.2. Emissions and boundary conditions
- 171 For the EU domain, the recently updated anthropogenic emissions for the year 2009
- 172 (http://www.gmes-atmosphere.eu/; Kuenen et al., 2014; Pouliot et al., 2014) were applied by
- all modelling groups and are based on the TNO-MACC-II (Netherlands Organization for
- 174 Applied Scientific Research, Monitoring Atmospheric Composition and Climate Interim
- 175 Implementation) framework. Annual emissions of methane (CH₄), carbon monoxide (CO),
- ammonia (NH₃), total non-methane volatile organic compounds (NMVOC), nitrogen oxides

 (NO_X) , particulate matter $(PM_{10}, PM_{2.5})$ and sulfur dioxide (SO_2) from ten activity sectors are 177 provided on a latitude/longitude grid of $1/8^{\circ} \times 1/16^{\circ}$ resolution. Emission inventories for the 178 NA domain were provided by US EPA and Environment Canada. The 2008 National 179 Emission Inventory (http://www.epa.gov/ttn/chief/net/2008inventory.html) and the 2008 180 Emission Modeling Platform (http://www.epa.gov/ttn/chief/emch/index.html#2008) with year 181 specific updates for 2006 and 2010 were used for the US portion of the modeling domain. 182 183 Canadian emissions were derived from the Canadian National Pollutant Release Inventory (http://www.ec.gc.ca/inrp-npri/) and Air Pollutant Emissions Inventory 184 (http://www.ec.gc.ca/inrp-npri/donnees-data/ap/index.cfm?lang=En) values for the year 185 2006. These included updated spatial allocations for Canadian mobile emissions (Zhang et al, 186 2012) for the emissions of NH₃ (Makar et al, 2009), as well as other updates (Sassi et al, 187 2010). Mexican emissions were 2008 projected forward from a 1999 inventory (Wolf et al, 188 2009). Seven pollutants (CO, NO_x, NH₃, SO₂, PM₁₀ PM_{2.5} and VOC) were used to develop 189 the model ready emission inventory. Further details and analyses of the anthropogenic 190 191 emissions used in both domains are provided in Pouliot et al. (2014). Annually-integrated anthropogenic emissions for both domains are presented in Table 2 while the spatial 192 distribution of NO_x emissions for the EU and NA domains are depicted in Fig. 1. Table 2 193 shows that anthropogenic emissions per km^2 in EU are larger than those in NA, except for 194 PM₁₀. Particularly NO_x and NH₃ emissions in EU are more than a factor of two larger than 195 196 those in NA. Consistent temporal profiles (diurnal, day-of-week, seasonal) and vertical distributions were also made available to maintain consistency among different groups. 197 NMVOC speciation factors were applied by all groups individually with a recommendation to 198 follow the NMVOC speciation profiles for EU by Visschedijk et al. (2007). The temporal 199 200 profiles for the EU anthropogenic emissions were provided from Schaap et al. (2005). Chemical and temporal profiles for the EPA anthropogenic emissions were based on the 201 2007v5 modeling platform (http://www.epa.gov/ttn/chief/emch/index.html#2008). 202 203 Each modeling group used their own biogenic emission module as detailed in Table 1. The majority of the models used the online MEGAN2 model (Model of Emissions of Gases and 204 Aerosols from Nature version 2; Guenther et al., 2006), two groups used the BEIS v3.14 205 model (Biogenic Emission Inventory System; Schewede et al., 2005) and one group (NL2) 206 used the Beltman et al. (2013) biogenic model. It should be noted that UK4 group used the 207 off-line simulated biogenic emissions provided by the Beltman et al. (2013) model. In 208 209 addition to the biogenic emissions algorithm used in the models, they may also differ in the 210 databases used for vegetation. Feedbacks may have a significant influence on biogenic emissions; reductions in biogenic isoprene emissions of 20% were found with the introduction 211 of the aerosol indirect effect (Makar et al., 2014a). The biogenic isoprene emissions 212 calculated on-line by each group show a large variability as shown in Table 2 that may lead to 213 large differences in the simulated O₃ levels. Curci et al. (2009) showed that different biogenic 214 emission models may lead to a factor of 2 difference in domain-integrated isoprene emissions 215

- over Europe while difference can be up to a factor of 5-6 locally. They estimated that these
- 217 differences on average may lead to an increase of 2.5 ppb in domain-mean surface O_3 levels
- and up to 10-15 ppb locally in the Mediterranean. Hourly biomass burning emissions were
- 219 provided by the Finnish Meteorological Institute (FMI) fire assimilation system

- 220 (<u>http://is4fires.fmi.fi/;</u> Sofiev et al., 2009). More details on the fire emissions and their
- uncertainties are discussed in Soares et al. (2014). The fire assimilation system provides only
- data for total PM emissions. Emissions of other species (CO, NO, NH₃, SO₂, NMVOC) were
- therefore deduced based on mass ratios relative to PM following Andreae and Merlet (2001).
- NMVOC speciation followed Wiedinmeyer et al. (2011) combined with the mapping to
- different chemical mechanisms proposed by Emmons et al. (2010). Note that the ES2a model
- does not include biomass burning emissions and as it does not contain aerosols leading to a
- 227 lack of effect of aerosols on photolysis rate calculations and therefore producing
- overestimated O_3 within the fire plumes (Badia and Jorba, 2014). Lightning NO_x is included
- in the UK4 model (O'Connor et al., 2014) as well as in the global MACC model used for the
- 230 boundary conditions as described below.
- 231 3-D daily chemical boundary conditions were taken from the MACC re-analysis (Inness et al,
- 232 2013). The MACC re-analysis (referred to as MACC hereafter) has been produced by
- assimilating satellite observations of O_3 , CO and NO_2 in the coupled system IFS-MOZART
- (Flemming et al., 2009). As pointed out in Inness et al. (2013), the assimilation of satellite-
- 235 corrected O_3 greatly improved the ozone total columns and stratospheric profiles but did not
- change significantly the surface levels because of the limited signal from this region in the
- assimilated satellite observations. The chemical species available in the reanalysis included
- O_3 , NO_x , CO, CH_4 , SO_2 , NMVOCs, sea-salt, dust, organic matter, black carbon and sulfate.
- 239 NMVOC species had to be lumped or disaggregated according to the individual models'
- 240 chemical speciation and particulate matter size discretization.

241 2.3. Observations

Measurements of hourly surface O₃ concentrations for the year 2010 in EU were taken from 242 the European Monitoring and Evaluation Programme (EMEP; http://www.emep.int/) and the 243 European Air Quality Database (AirBase; http://acm.eionet.europa.eu/databases/airbase/) and 244 in NA from the Canadian National Atmospheric Chemistry (NAtChem) Database and 245 Analysis Facility operated by Environment Canada (http://www.ec.gc.ca/natchem/) that 246 247 contains measurements from the Canadian National Air Pollution Surveillance Network (http://maps-cartes.ec.gc.ca/rnspa-naps/data.aspx), the Canadian Air and Precipitation 248 Monitoring Network (http://www.ec.gc.ca/natchem/), the U.S. Clean Air Status and Trends 249 Network (http://java.epa.gov/castnet/clearsession.do), the U.S. Interagency Monitoring of 250 Protected Visual Environments Network (http://views.cira.colostate.edu/web/DataWizard/), 251 and the U.S. Environmental Protection Agency's Air Quality System database for U.S. air 252 quality data (http://www.epa.gov/ttn/airs/airsaqs/detaildata/downloadaqsdata.htm). In the 253 AQMEII2, rural, urban and suburban background stations were extracted from the EMEP and 254 AirBase networks. Given the coarse native grid resolutions used in different models (Table 1), 255 data from only rural background stations was used in the comparisons. Stations that have 256 more than 90% data availability have been selected for the comparisons. Regarding the whole 257 simulation domains, hourly surface O₃ observations were provided by 510 and 200 stations in 258 EU and NA, respectively. A geographical break-down into four sub-regions for each 259 260 continent has also been defined based on the climatological and source characteristics. The geographical break-down of these stations overlaid with the annually-averaged anthropogenic 261

NO_x emissions is shown in Fig.1. Model evaluation statistics were computed for the four sub-262 regions separately. The European sub-region EU1 is characterized by north-western European 263 sources with a transition climate between marine and continental and hosts 102 stations. Sub-264 region EU2 covers the north-eastern and central Europe sources as well as Germany with 277 265 monitoring stations. Sub-regions EU3 and EU4 are characterized by a Mediterranean type 266 climate. Sub-region 3 covers south-western sources including Italy (30 stations) while sub-267 268 region 4 covers the East Mediterranean with 101 stations. The North American sub-region 1 (NA1) covers the western U.S. and south western Canada with 80 stations. It includes large 269 emission sources along the coast as well as polluted hot spots like Los Angeles that are 270 characterized by poor air quality. NA2 consists of U.S. plains and covers 36 monitoring 271 stations and is characterized by a continental and humid climate. NA3 consists of north 272 eastern NA and south central Canada and is characterized by the largest emissions in North 273 America and contains 60 monitoring stations. Finally NA4 covers the south eastern part of 274

275 U.S., consisting of 24 monitoring stations.

276 To evaluate the capability of the modeling systems to simulate the tropospheric distribution of

277 O₃ concentrations, comparisons against O₃ soundings provided by the World Ozone and

278 Ultraviolet Radiation Data Centre (WOUDC: <u>http://www.woudc.org/</u>) have been carried out.

279 Ozone concentration data from nine stations in EU and six stations in NA have been used for

the comparisons. For an optimal comparison with observations, model profiles were

computed by averaging only over the available observation hours. The participants were

required to provide their data at fixed heights up to 18 km above the ground in order to be

- comparable. However, due to the coarse vertical resolution of some models in the upper
- troposphere and not simulating the stratospheric chemistry, the analyses are performed only

for the first 9 km above ground.

286 2.4. Statistical analyses

To score the individual model performances as well as those of the ensemble mean and median, the following statistical parameters have been calculated: Pearson's correlation coefficient (*PCC*: Eq.1), root mean square error (*RMSE*: Eq.2); normalized mean standard error (*NMSE*: Eq.3) and normalized mean bias (*NMB*: Eq.4).

291
$$PCC = \left[\frac{\frac{1}{N}\sum_{i=1}^{N} (O_i - \overline{O})(P_i - \overline{P})}{\sigma_o \sigma_P}\right]$$
(Eq. 1)

292
$$RMSE = \sqrt{\frac{1}{N} \sum_{i=1}^{N} (P_i - O_i)^2}$$
 (Eq. 2)

293
$$NMSE = \frac{\sum_{i=1}^{N} (P_i - O_i)^2}{N \times \overline{P} \times \overline{O}} \times 100$$
(Eq. 3)

$$NMB = \frac{\sum_{i=1}^{N} (P_i - O_i)}{\sum_{i=1}^{N} O_i} \times 100$$

i=1

(Eq. 4)

where P and O denote model predictions and observations, respectively. The PCC is a 295 measure of associativity and allows gauging whether trends are captured, and it is not 296 sensitive to bias; *RMSE* is a measure of accuracy and, because it is squared, is sensitive to 297 large departures. NMSE and NMB are normalised operators, useful for comparing scores 298 coming from time series of different lengths, as those produced over different areas and/or 299 with different time span. The comparison is performed individually for the two domains and 300 their sub-regions for the whole year of 2010 and on a seasonal basis, in order to identify 301 302 which regions and/or seasons lead to systematic errors.

303

304 3. Results and Discussion

305 3.1. Surface ozone analyses

Observed and simulated diurnal cycles of surface O₃ concentrations averaged over the whole
simulation period (2010) are shown in Fig. 2a,d for EU and NA, respectively. Models are
labelled by the ID of the respective modeling group, with each ID corresponding to a member
of the overall model ensemble. In the same figures, the MACC IFS-MOZART global model
(MACC) and the ensemble mean and median are also shown. Note that the MACC model is

- 311 not considered in the ensemble calculations.
- 312 *3.1.1. Europe*

Most models capture reasonably well the shape of the annual diurnal cycle over Europe as 313 seen in Fig.2. The temporal variations on all time scales were captured successfully as seen in 314 Table 3 (*PCC*>0.80), although the predicted O_3 levels are generally underestimated by up to 315 18%. Only one group (UK4) slightly overestimates the yearly-averaged observed surface O₃ 316 317 levels by 2% while the other groups have underestimations up to 18%. The largest underestimations are calculated for IT2 (by 16%) and CH1 (by 18%) groups. Other groups 318 have mean normalized biases within the $\pm 5\%$ to 15% range suggested by Russell and Dennis 319 (2000). Fig.2a shows that the underestimations generally occur both during day and night 320 hours, which is expected to some extent given the coarse horizontal resolution (Qian et al., 321 322 2010). The exceptions are AT1, DE4, SI1 and UK4 that overestimate the night time levels. The MACC model underestimates the nighttime levels as also reported in Innes et al. (2013). 323 Overestimation of nighttime O₃ levels can be due to the overestimation of NO₂ concentrations 324 under low-NO_x conditions leading to overestimated O₃ concentrations (e.g. DE4). Fig.2a 325 326 shows that the underestimations generally occur both during day and night hours, which is expected to some extent given the coarse horizontal resolution (Qian et al., 2010). The 327 exceptions are AT1, DE3, DE4, SI1, and UK4 that overestimate the night time levels. The 328 MACC model underestimates the nighttime levels as also reported in Innes et al. (2013). The 329

small overestimation of nighttime O₃ levels for AT1 and SI1 can be attributed to the 330 underestimation of nocturnal ozone titration in urban areas with high NO_x emissions for the 331 QSSA solver that was applied for these simulations. For DE4, where a modified version of 332 this solver (Forkel et al., 2014) has been applied, the overestimation of nighttime ozone can 333 be attributed to a general overestimation of NO₂ concentrations under low-NO_x conditions. 334 This is also the case for the DE3 model during the nighttime, where this overestimation is 335 336 probably related to difficulties of the meteorological model to simulate nighttime vertical mixing accurately and, furthermore, to comparatively small dry deposition fluxes for O₃ 337 simulated by the model (see Table 3). It should be noted that the ES2a model does not include 338 anthropogenic aerosols and secondary aerosol formation and neither aqueous chemistry, 339 leading to a more oxidized atmosphere. Furthermore, the heterogeneous formation of HNO₃ 340 through N₂O₅ hydrolysis, which is an important sink of NO₂ during night, is not considered in 341 ES2a (Badia and Jorba, 2014). As a consequence, the ES2a model overestimated the annual 342 domain-mean NO₂ levels by 15% while the rest of the models underestimate NO2 by 9% to 343 344 45%. The overestimation of surface O₃ levels by the ES2a model can also partly be due to the coarser vertical resolution of its first layer (45 m) compared to other models (Table 1). The 345 general underestimation may be partly attributed to biases in meteorological variables, 346 including an overestimation of surface wind speeds by all models by up to 60% and a general 347 slight underestimation of surface temperatures by less than 1 K (Brunner et al., 2014). Such a 348 small temperature bias, however, will affect ozone levels by no more than a few ppb (Sillman 349 and Samson, 1995). A common feature of all groups is that the daily maximum is simulated 350 earlier than the observed maximum. Differences in O₃ predictions between the WRF-CHEM 351 models suggest that the choice of the chemical mechanism plays an important role in the 352 353 model performance. WRF-CHEM runs using RADM chemical mechanism (AT1, ES1 and SI1) produced higher concentrations than runs using RACM (IT2) and CBMZ (ES3 and IT1) 354 mechanisms (Baro et al., 2014). These differences may partly be attributed to VOC emission 355 preprocessing. WRF-CHEM is designed to ingest VOC emissions for RADM2 and then, in 356 357 case of other mechanisms, the emissions are chemically specified to the final scheme, possibly leading to a degradation of the reactivity in the VOC mixture. There are also 358 differences in the microphysics schemes among the different WRF-CHEM configurations 359 used, leading to different cloudiness and therefore to different temperature and radiation 360 acting on the O₃ production (Brunner et al., 2014; Baro et al., 2014). Makar et al. (2014a) and 361 Wang et al. (2014) found that models including the simulation of indirect effects tended to 362 have lower O₃ concentrations during the summer production period than those with the direct 363 effect only, or those with no feedbacks. This is due to the reduction of NO₂ mixing ratios 364 during daytime and near-surface temperatures, resulting from the reduction of solar radiation 365 (Wang et al., 2014). Dry deposition of O₃ is also investigated for the models that provided 366 deposition data (CH1, DE3, DE4, ES1, ES2a, ES3, IT2, NL2 and SI1) in order to explain the 367 differences in simulated O₃ levels among the models (Table 3). The results show a negative 368 relation between underestimation and dry deposition; i.e. the underestimation increases with 369 370 decreasing deposition, suggesting that other terms aside from deposition were controlling the O₃ concentrations (chemistry, vertical diffusion etc.). 371

The model performances are also assessed against the observed variability in box-and-whisker 372 plots of Fig.2b and e. The plot shows the frequency distribution of observed and simulated 373 surface O₃ mixing ratios. The spread of the data in the European case is largest in CH1, ES2a 374 and UK4 (Fig. 2b). The majority of other models show a much lower spread, which also tends 375 to be lower than the observed spread. Data from MACC are associated with a larger spread 376 compared to the observations in both domains, suggesting a better representation of local 377 378 processes by regional models as well as an indication of an exaggerated seasonal cycle simulated by the MACC model. The larger spread in some models as compared to others is 379 partially related to the amplitude of the diurnal ozone cycle, which tends to be larger in 380 models simulating a more stable and shallow nocturnal PBL such as the global MACC model 381 (Innes et al., 2013). A larger amplitude may also be expected for models with a higher vertical 382 resolution. The NMB vs NMSE plot (also known as the soccer diagram) for EU (Fig. 2c) 383 shows that the models have mean biases below 30% and mostly below 15%. The geographical 384 analyses for the EU domain presented in Fig.3 show that for the majority of models, the 385 386 underestimation is mainly originating from sub-region EU2 (north Eastern Europe) while in sub-region EU4 (East Mediterranean), most models overestimate the observed mean. The 387 underestimation, particularly in EU1 and EU2 could be partly due to the chemical boundary 388 conditions (Fig.3) as discussed in more detail in Sect. 3.3. 389

390 3.1.2. North America

The hourly O₃ temporal variability over the whole simulation period is also well captured 391 (PCC>0.78) by all groups for the NA domain (Table 3). The CA2f model overestimates the 392 nighttime surface O₃ concentrations and underestimates the daytime levels with a slight 393 overall overestimation of 2% while other groups underestimate the nighttime levels (Fig.2d). 394 NMSE values are below 10% for all the groups while NMB values are within $\pm 15\%$ except for 395 the US8 model, which underestimates the surface O₃ levels by 22%. The box plots for the NA 396 case (Fig. 2e) shows that the MACC model has the highest variability while CA2f is 397 398 characterized with the smallest spread. Larger biases in US7 and US8 can also be partly attributed to their coarser resolution (36 km) compared to other NA models (Table 1). In the 399 NA case, according to the soccer diagrams (Fig. 2f), all groups and sub-regions are 400 characterized with biases lower than 25% except for US8. The geographical break down 401 402 presented in Fig.4 shows that the US8 model underestimates in all sub-regions. The MACC model also shows a general underestimation in all sub-regions except for NA4. Regarding the 403 dry deposition of O_3 (Table 3), the results suggest that the large underestimation by US8 can 404 be partly due to the relatively large O_3 dry deposition simulated by the model, acting as a 405 406 significant sink. As analyzed in Yahya et al. (2014a,b) and Wang et al. (2014), other factors that contribute to underpredictions of O₃ by the US8 model include large underpredictions of 407 afternoon temperatures, low MACC boundary conditions of O₃, the overpredictions of the 408 NO_x titration effects on O₃ during nighttime, possible underestimates in biogenic VOCs and 409 wildfire emissions, and the inclusion of aerosol indirect effects. The lower spread in CA2f 410

- 411 seems to be due to overpredicting the lower end of the O_3 range compared to the observations, 412 in majors NA2 and NA4
- 412 in regions NA3 and NA4.
- 413 3.2. Seasonal vs. geographical surface ozone variations

414 *3.2.1. Europe*

415 Inter-seasonal variations of surface O_3 concentrations are analyzed for each sub-region in 416 order to understand how the model bias varies depending on the region and season. The 417 results for the EU domain are depicted in Fig. 5. The temporal variability in Europe is better

- 418 captured in all models in summer and autumn (PCC=0.8-0.9) than in winter and spring
- 419 (PCC=0.6-0.8). There is a systematic overestimation of the observed concentrations in
- 420 autumn by up to 35%, particularly by the DE4 model. In winter (Fig. 5a), O_3 mixing ratios in
- 421 EU2 are underestimated by more than 50% by three groups (CH1, ES2a and UK4), which 422 also underestimate systematically in other sub-regions, probably due to the bias from the
- boundary conditions from the MACC model. The MACC model underestimates by largest
- 424 during winter (by 8% to 55%) and overestimates by largest in autumn (by 8% to 25%).
- 425 Regarding EU1, all groups are within the 30% bias range. Spring and summer O_3 mixing
- ratios (Fig. 5b,c) in all EU sub-regions are similarly reproduced by all groups, with error
- below 30%. In autumn, the majority of the models are biased high. In northern Europe (EU1
- 428 and EU2), the majority of the models underestimate O_3 levels in all seasons with the DE4,
- 429 UK4, and ES2a models overestimating during summer. There is a general overestimation in
- autumn in the EU1 sub-region by all models except for CH1 and IT2. The models NL2, DE4,
- 431 UK4 and ES2a overestimate the summertime O_3 levels in southern Europe. The East
- 432 Mediterranean region (EU4) is characterized by overestimated O_3 levels, in particular during
- autumn. The results show that the largest underestimations were calculated for the EU2
- region, which is characterized with large anthropogenic emissions in the Eastern Europe that
- 435 may lead to overestimated O_3 -titration by NO_x .

436 *3.2.2. North America*

- Inter-seasonal and geographical variations of the models performances in NA are presented in
 Fig.6. US8 underestimates the observations in all seasons and in particular in winter and
- spring, and much larger compared to other models. In sub-region NA1, US6 overestimates by
- 440 up to 9% while US8 underestimates by up to 22% in all seasons. CA2f slightly overestimates 441 the winter and autumn O_3 levels by 3% and 5%, respectively. In the sub-regions NA2 and
- the winter and autumn O_3 levels by 3% and 5%, respectively. In the sub-regions NA2 a NA3, there is a general underestimation of all O_3 in winter and spring and a general
- 443 overestimation in summer and autumn except for the US8 model. The winter and spring
- 444 underestimates may be the result of underpredictions of afternoon temperatures and excessive
- 445 O_3 titration by NO_x as NA3 can be characterized by the largest emission sources in NA. In
- 446 NA4, summertime O_3 levels are overestimated by all models including the US8 model.
- 447 Slightly lower correlation coefficients (*PCC*=0.7-0.9) are calculated for winter in NA while
- other seasons are simulated with *PCC* values of ~0.8-0.9, with slightly lower PCC values
- 449 calculated for US7 (not shown).
- 450 3.3. Influence of chemical boundary conditions
- 451 The influence of the chemical boundary conditions on the simulated surface O_3 levels has also
- been investigated on a seasonal basis. The analysis is carried out for the EU2 (north Eastern
- Europe) sub-region for Europe assuming that it is the least affected by the dominant westerly
- transport and having large anthropogenic emissions, suggesting that O_3 levels are more

strongly controlled by local processes than regional transport, compared to the other sub-455 regions. Following the same rationale, sub-region NA3 was selected for the NA domain. The 456 results presented in Fig.7a show that in winter, all models underestimate O₃ levels along with 457 458 the MACC model that provides the boundary conditions suggesting that large scale circulation and chemistry dominates over the local O₃ production. In spring and in summer 459 (Fig.7b,c), the regional production is more important than transport due to increased 460 461 photochemical activity. In autumn (Fig.7d), transport becomes more effective over local production. The MACC model slightly overestimates the summer levels (NMB=1%), and 462 slightly underestimates the autumn levels (NMB=-5%) while it underestimates the winter and 463 spring levels 55% and 21%, possibly leading to the systematic overestimation of the regional 464 models in autumn. The impact of large-scale transport over NA is less pronounced compared 465 to Europe (Fig.8). The impact is the smallest during summer when photochemical production 466 is the largest (Fig.8c). At the same time, it is interesting to note that the MACC results in the 467 winter for NA1 are the lowest of the models shown in Fig.8a, with a deficit of 8 ppb relative 468 to the observations at 0 LST. The implication is that local chemistry, physics, model 469 resolution and/or emissions relative to the global model all account for an increase in the 470 winter O_3 levels for region NA1 of 8 ppb (28.5%), and these local effects are captured by the 471 suite of regional models. This may be compared to findings from the HTAP experiment, 472 which suggest a 20% reduction in emissions in Europe, South Asia and East Asia would result 473 474 in a 0.9 ppb reduction in O_3 in North America (Reidmiller et al, 2009). Here, simulated O_3 levels seem to be much more sensitive to the local O_3 chemistry than to the boundary 475 conditions associated with long-range transport (winter being the dominant season for long-476 range transport effects). Over both continents, the nighttime differences in all seasons are 477 478 particularly large, with the MACC model largely underestimating the nighttime O₃. Similar results were reported by Solazzo et al. (2012 and 2013a) for the first phase of the AQMEII 479 project. A more detailed analysis of the influence of the MACC boundary conditions on a 480 range of simulated species is presented in Giordano et al. (2014). 481

482 3.4. Multi-model mean and median

The combination of concentrations simulated by several models can enhance the skill when 483 compared to those from individual models (Galmarini et al., 2004a,b), which has also been 484 485 demonstrated by Solazzo et al. (2012) in the first phase of the AQMEII project. In the present study, we provide simple multi-model mean and median analyses. Therefore, the calculated 486 multi-model mean and median presented in Table 3-5 and in Fig.2-11 can only provide a 487 basic distribution of all models with respect to the observations and should not be treated as 488 multi-model ensemble analyses as they represent the bias originating from each individual 489 model. As shown in Solazzo et al. (2012, 2013b) and Kioutsioukis and Galmarini (2014), 490 introducing correlated biases into ensembles and analysis of the redundancy of the datasets is 491 essential. As detailed multi-model ensemble analysis is not the scope of this paper, further 492 analyses have been performed by Kioutsioukis et al. (2014) for the EU case using the multi-493 494 model data presented in the present paper.

495 3.5. Regulatory analysis based on 8-hour maximum surface O₃

- 496 Observed and simulated daily maximum 8-hour averaged surface O_3 levels during the O_3
- 497 season (May-September), which is a regulatory metric used in EU and NA, are compared in 498 order to understand how the model biases vary with O_3 levels. The results are shown in Fig.9
- order to understand how the model biases vary with O_3 levels. The results are shown in Fig (note that in Fig.9, observed concentrations are presented by /10). Over EU, all models
- 500 overestimate O_3 concentrations below 50 µg m⁻³ by ~40% to ~80% while they underestimate
- values above 140 μ g m⁻³ except for the UK4 model that overestimates the levels above 160 μ g
- 502 m^{-3} . Most models follow the MACC model up to a concentration of 200 µg m^{-3} with
- 503 increasing variability towards higher concentrations. NL2 and UK4 models overestimate the
- 504 $230-240 \,\mu \text{g m}^{-3}$ concentration bin where the spread is also largest among other models. The
- 505 UK4 model defines the upper boundary while IT2 defines the lower boundary of the envelope
- until 100 μ g m⁻³ while above that, the highest differences are calculated for IT1. The CH1
- model, which together with the IT2 model showed the largest negative biases in annual mean
 values, is more consistent with other models when considering 8-hour maximum values.
- Above a concentration of 70 μ g m⁻³, ES2a, NL2 and UK4 models are associated with positive
- 510 deviations from the MACC model while other models are below the MACC-simulated levels.
- 511 Results show that depending on the station, there are underestimations by up to $>200 \ \mu g \ m^{-3}$.
- 512 Over NA (Fig.9b), the biases are lower compared to EU. Note that for NA the values are
- reported in volume mixing ratios (ppb) rather than concentrations ($\mu g/m^3$). The surface O₃
- levels below 30 ppb are overestimated by all models by ~15-25% and levels above 60 ppb are
- underestimated by all models by up to ~80%. The largest biases are calculated for US8 except
- 516 for the 120-130 ppb bin where US7 has the largest bias. US8 has the smallest bias below 50
- 517 ppb. The results show that models have a tendency to severely underpredict high O_3 values
- 518 which are of concern for air quality forecast and control policy applications. Further
- improvement of model treatments (e.g., gas-phase chemistry, O_3 dry deposition and processes
- 520 affecting afternoon temperature predictions) and inputs (e.g., boundary conditions, biogenic
- 521 VOCs and wildfire emissions) as well as a better understanding of interplays among on-line-
- 522 coupled atmospheric processes (e.g., the impact of aerosol indirect effects on O_3 formation)
- 523 are urgently needed.
- 524 3.6. Vertical ozone profiles

The model results from each group as well as the ensemble mean and median are compared 525 with O₃ soundings obtained from WOUDC for the EU and NA domains up to 9 km height 526 above the ground. Figs.10 and 11 show the observed and simulated vertical O₃ levels at fixed 527 heights over the EU and NA domains, respectively while Tables 4 and 5 present the 528 normalized mean bias (NMB) for all the models and ensemble mean and median. On average, 529 most models underestimate the observed vertical profiles by up to 22% over EU. The DE4 530 531 model generally has smaller biases compared to other groups except for the station STN156 where it overestimates by $\sim 12\%$ (Fig.10). The ensemble mean/median improves the results 532 compared to the majority of the models depending on the station. The ensemble mean results 533 in smaller biases compared to the median. Over NA (Fig.11), the CA2f model underestimates 534 535 the vertical O₃ levels at all stations by10-17% (Table 5). US6 and US7 have the smallest biases in most stations but with overestimations of 14% and 5%, respectively, at STN457. The 536

537 US8 model underestimates at all stations by 4-15% but overestimates at STN457 by 2%. The

- ensemble mean and median lead to improved results compared to CA2f at all stations above
- \sim 1000-2000 m and to US8 at STN107 and STN456 below 2000-3000 m. Over Europe, among
- 540 others, STN318 station (Valentia Observatory, Ireland) can be considered as a site that is
- 541 largely impacted by long-range transport and is associated with the largest underestimation 542 (NMB = -11%) by the MACC model (not shown), suggesting that boundary conditions can
- (NMB = -11%) by the MACC model (not shown), suggesting that boundary conditions can partly contribute to the underestimated vertical profiles by a majority of the models. Results
- also show that the tropospheric biases in the MACC model (Fig.10,11) are less pronounced
- than the surface bias as also shown by Inness et al. (2013).
- 546

547 4. Summary and Conclusions

548 An operational evaluation of simulated ozone (O₃) levels over Europe (EU) and North

549 America (NA) in 2010 using eight different on-line-coupled air quality models from sixteen

- groups has been conducted in the context of the AQMEII project. Seven groups from EU and
- two groups from NA applied the WRF-CHEM model, but with different settings.

552 Anthropogenic emissions and chemical boundary conditions were prescribed while biogenic

emissions were calculated online by each individual group. All groups interpolated their

- model output to a common output grid and a common set of receptor locations and uploaded
- the data to the ENSEMBLE system. The results are evaluated against surface and sounding observations, which are provided by operational over EU and NA, at continental and sub-
- 557 regional levels on annual and seasonal basis.

All models capture, reasonably well, the shape of the domain-averaged annual diurnal cycle 558 of O₃ over both domains, while the sub-regional temporal variability are simulated from 559 moderate to good depending on the season and the sub-region that the particular model is 560 configured for. There is a general underestimation of the annual surface O₃ by up to 18% and 561 22% over EU and NA, respectively. Differences in performance among models can be 562 attributed partly to the chemical mechanism used in the models, partly to VOC preprocessing 563 and different biogenic emissions, and partly to the differences in the microphysics, leading to 564 different cloudiness and therefore to different photolysis, temperature and radiation acting on 565 the O_3 production. The sub-regional analyses highlight the influence of the anthropogenic 566 emissions while the seasonal analyses show a strong tendency to overestimate the autumn 567 surface levels. The temporal variation and magnitudes are much better captured during 568 summer compared to other seasons. The winter and spring underestimations may be resulting 569 570 from underprediction of afternoon temperatures, excessive O_3 titration by too much NO_x as well as biases from the chemical boundary conditions. Boundary condition analyses show that 571 wintertime levels are mostly driven by transport rather than local production due to limited 572 photochemistry. The global MACC model providing the boundary conditions to the regional 573 574 models largely underestimate the surface ozone levels particularly in winter, leading to a 575 negative bias in the regional model simulations, while in most sub-regions, it largely

- 576 overestimates the autumn O_3 levels in winter, leading to the systematic overestimations of
- surface autumn O_3 levels by the regional models. The inclusion of aerosol indirect effects in

- some online-coupled models also contributes in part to the underpredictions of O₃ mixing 578
- ratios. On average, most models underestimate the observed vertical profiles by up to 22% 579
- over EU and up to 17% over NA. 580
- Comparison of observed and simulated daily maximum 8-hour averaged surface O₃ levels 581
- during the O₃ season (May-September), which is a regulatory metric used in EU and NA, 582
- show that over Europe, O_3 concentrations below 50 µg m⁻³ are overestimated by up to 80% 583
- while levels above 140 μ g m⁻³ are underestimated. Over NA the surface O₃ levels below 30 584
- ppb are overestimated by all models by up to 25% and levels above 60 ppb are 585
- underestimated by all models by up to 80%. This has implications for air quality forecast and 586
- policy applications. 587
- Overall, the results show a slight improvement in the surface ozone level predictions over EU 588
- by the models that participated in the second phase of AQMEII compared to those that 589
- participated in the first phase. The NMB calculated for the whole domain and simulation 590
- period in the first phase ranged from -24% to 9% while in this second phase, the NMB range 591
- was calculated to be -18% to 2%. On the other hand over NA, there is a significant change 592
- 593 between the two phases of the project: the overestimation of 3% to 22% in the first phase
- shifted to a NMB range of -22% to 3%. These results, however, should not be considered as 594
- solely the difference between on-line and off-line models as different simulation years, 595 different emissions, different sets of models, particularly for the NA case, and different
- 596 boundary condition data should be taken into account. Additionally, as the results presented in
- 597
- this paper are temporally and spatially averaged, cases where feedback mechanisms are of 598
- importance must be further studied and evaluated. 599
- 600
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- extracted from NAtChem's precipitation-chemistry data base and were provided by several 613
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- RObotic NETwork (AeroNet) and its data-contributing agencies provided North American 617

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	Groups	Domain	Model	Grid Spacing	First layer height (m)	Biogenic Model	Gas Phase	Photolysis	Model Reference
M1	AT1	EU	WRF-CHEM	23 km	24	MEGAN	RADM2 ¹	Fast-J ¹³	Grell et al., 2005
M2	CH1	EU	COSMO-ART	0.22°	20	Gunter et al., 1998	RADM2K ²	GRAALS+STAR ²	Vogel et al., 2009
M3	DE3	EU	COSMO-MUSCAT	0.25°	20	Gunther et al., 1993	RACM-MIM2 ³	Fast-J	Wolke et al., 2012
M4	DE4	EU	WRF-CHEM	23 km	24	MEGAN	RADM2 modified ⁴	Fast-J	Grell et al., 2005; Forkel et al., 2014
M5	ES1	EU	WRF-CHEM	23 km	24	MEGAN	RADM2	Fast-J	Grell et al., 2005
M6	ES2a	EU	NMMB-BSC-CTM	0.20 °	45	MEGAN	CB05 ⁵	Fast-J	Jorba et al., 2012
M7	ES3	EU	WRF-CHEM	23 km	24	MEGAN	CBMZ ⁶	Fast-J	Grell et al., 2005
M8	IT1	EU	WRF-CHEM	23 km	24	MEGAN	CBMZ	Fast-J	Grell et al., 2005
M9	IT2	EU	WRF-CHEM	23 km	24	MEGAN	RACM ⁷ Fast-J		Grell et al., 2005
M10	NL2	EU	RACMO LOTOS-EUROS	0.5 ° ×0.25°	25	Beltman et al., 2013	CB-IV modified ⁸	Poppe et al., 1996	Sauter et al., 2012
M11	SI1	EU	WRF-CHEM	23 km	25	MEGAN	RADM2	Fast-J	Grell et al., 2005
M12	UK4	EU	MetUM-UKCA RAQ	0.22 °	20	TNO	UKCA RAQ ⁹	Fast-J	Savage et al., 2013
M13	CA2f	NA	GEM-MACH	15 km	20.66	BEIS	ADOM-II ¹⁰	Dave, 1972	Makar et al., 2014a,b
M14	US6	NA	WRF-CMAQ	12 km	19	BEIS3.14	CB05-TU ¹¹	Binkowski et al., 2007	Wong et al., 2012
M15	US7	NA	WRF-CHEM	36 km	55-60	MEGAN	MOZART ¹²	fTUV ¹⁴	Grell et al., 2005
M16	US8	NA	WRF-CHEM	36 km	38	MEGAN	CB05	fTUV	Grell et al., 2005; Wang et al., 2014

Table 1. Modelling systems participated to AQMEII2 and their configurations

1 Stockwell et al. (1990); 2 Vogel et al. (2009); 3 Karl et al. (2006); 4 Forkel et al. (2014); 5 Yarwood et al. (2005); 6 Zaveri et al. (1999); 7 Stockwell et al. (1997); 8 Sauter et al. (2012); 9 Savage et al. (2013); 10 Lurmann et al. (1986); 11 Whitten et al., 2010; Sarwar et al., 2011; 12 Emmons et al. (2010); Knote et al. (2013); 13 Wild et al., 2000; 14 Tie et al., 2003

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Table 2. Annual anthropogenic emissions (ktons km⁻² yr⁻¹) provided by TNO-MACC-II inventory and biogenic isoprene emissions (ktons km⁻² yr⁻¹) integrated over the EU and NA domains.

Species	EU	NA
CO	614	478
NO _x *	277	120
NMVOC	230	85
NH ₃	109	31
SO ₂	109	70
PM _{2.5}	49	29
PM ₁₀	69	76
ISOP ^{**}	2.4-24.9	0.02-8.1

* Only anthropogenic NOx is reported.

**The groups that provided isoprene emissions are AT1, CH1, DE3, IT2, NL2 and UK4 for the EU domain and CA2f, US6 and US7 for the NA domain.

Members	r	NMSE (%)	NMB (%)	<i>RMSE</i> [*]	Dry Deposition (Tg km^{-2})
M1/AT1	0.86	2.66	-4.92	9.57	NP
M2/CH1	0.82	8.03	-18.30	15.42	0.28
M3/DE3	0.68	6.37	-2.12	15.02	0.13
M4/DE4	0.83	3.17	-1.64	10.62	2.24
M5/ES1	0.86	4.08	-11.41	11.44	2.18
M6/ES2a	0.83	6.37	-7.71	14.59	2.79
M7/ES3	0.86	4.29	-12.07	11.69	1.82
M8/IT1	0.85	4.57	-12.45	12.03	NP
M9/IT2	0.84	6.21	-15.80	13.76	1.77
M10/NL2	0.89	2.83	-4.34	9.90	0.14
M11/SI1	0.87	2.38	-3.78	9.10	1.91
M12/UK4	0.85	7.88	2.30	17.08	NP
EU Mean	0.86	3.22	-7.70	10,37	
EU Median	0.86	3.23	-8.69	10.33	
M13/CA2f	0.85	1.45	2.43	4.02	0.09
M14/US6	0.84	2.15	1.14	4.85	0.10
M15/US7	0.78	4.36	-4.56	6.72	0.15
M16/US8	0.88	8.11	-22.36	8.26	3.05
NA Mean	0.83	3.70	-11.98	5.94	
NA Median	0.87	2.62	-9.51	5.07	

Table 3. Statistical comparisons of observed and simulated annual domain-mean hourly surface O_3 and domain- and annually-integrated O_3 dry deposition over EU and NA in 2010.

* *RMSE* is in units of µg m⁻³ for EU and ppb for NA.

Stations	Station Name	Country	Lat/Lon	AT1	CH1	DE3	DE4	ES1	ES2a	ES3	IT1	IT2	NL2	SI1	UK4	Mean	Median
STN043	Lerwick	United Kingdom	60.1/-1.2	-8.40	-12.14	-27.80	-2.39	-9.82	-11.46	-7.86	-6.91	-9.05	-3.32	-8.16	-11.13	-10.40	-7.40
STN053	UCCLE	Belgium	50.8/4.4	-4.11	-10.09	-14.08	3.80	-6.02	-7.46	-4.14	-4.53	-7.23	-1.96	-3.95	-4.86	-5.50	-3.56
STN099	Hohenpeissenberg	Germany	47.8/11.0	-10.65	-21.94	-23.98	-2.04	-12.15	-11.98	-8.96	-9.55	-11.47	0.17	-10.39	-8.43	-11.62	-9.63
STN156	Payerne	Switzerland	46.5/6.6	1.18	-10.06	-11.71	11.77	-0.63	1.84	2.51	2.43	0.51	2.70	1.44	3.94	0.64	2.52
STN242	Praha	Czech Rep.	50.0/14.5	-8.55	-16.18	-26.48	-1.72	-11.38	-8.82	-6.98	-6.97	-8.68	-4.77	-7.86	-5.06	-9.50	-7.35
STN308	Barajas	Spain	40.5/-3.7	-6.02	-14.29	-9.83	1.91	-7.72	-4.77	-6.72	-6.32	-7.83	0.21	-5.67	-1.61	-5.95	-5.01
STN316	De Bilt	Netherlands	52.1/5.2	-4.57	-5.83	-9.82	3.62	-6.14	-4.29	-4.99	-5.08	-7.29	1.15	-4.37	-0.59	-4.23	-3.76
STN318	Valentia	Ireland	51.9/-10.3	-6.51	-10.56	-15.49	-0.44	-8.01	-9.30	-6.00	-2.93	-6.35	-5.74	-6.43	-5.97	-7.01	-5.04
STN348	Ankara	Turkey	40.0/32.9	-11.48	-16.13	-12.94	5.76	-13.38	-4.28	-10.95	-11.12	-15.24	0.55	-11.36	2.41	-8.66	-9.74

Table 4. *NMB* calculated for vertical O₃ profiles for each model group and ensemble mean and median for the WOUDC stations in EU.

Table 5. *NMB* calculated for vertical O₃ profiles for each model group and ensemble mean and median for the WOUDC stations in NA.

Stations	Station Name	Country	Lat/Lon	CA2f	US6	US7	US8	Mean	Median
STN021	Stony Plain	Canada	53.4/-114.1	-9.82	1.58	-2.29	-4.71	-3.81	-2.85
STN107	Wallops Island	USA	37.9/-75.5	-10.19	1.77	-1.17	-13.52	-5.78	-6.30
STN338	Bratts Lake	Canada	50.2/-104.8	-14.29	0.27	-3.26	-9.46	-6.68	-4.47
STN456	Egbert	Canada	44.2/-79.8	-16.78	-1.40	-3.95	-15.01	-9.28	-8.54
STN457	Kelowna	Canada	49.9/-119.4	-10.09	13.61	4.95	-0.62	1.96	2.05
STN458	Yarmouth	Canada	43.9/-66.1	-17.76	-1.17	-5.95	-15.27	-10.04	-10.20

Figure Captions

Fig.1. Annual NO_x emissions (tonnes/grid) overlaid with the rural monitoring stations used for model performance evaluation in EU (a) and in NA (b). The red circles show EU1/NA1, yellow diamonds show EU2/NA2, green squares show EU3/NA3 and black triangles show EU4/NA4.

Fig.2. Observed and simulated annual mean diurnal profiles (a,d), box plots (b,e) and soccer diagrams (c,f) for surface levels ozone mixing ratios in EU (upper panel) and NA (lower panel). Mn and Md represent the mean and median ensembles, respectively. EU0 and NA0 represent the two respective continents. Different colors represent the different sub-regions. Note the differences in scales.

Fig.3. Geographical distributions of observed and simulated annual surface level ozone mixing ratios in EU. Note the differences in scales.

Fig.4. Geographical distributions of observed and simulated annual surface level ozone mixing ratios in NA. Note the differences in scales.

Fig.5. Soccer diagrams for the seasonal and geographical model performances in EU: a) winter, b) spring, c) summer and d) autumn. Mn and Md represent the mean and median ensembles, respectively. EU0 and NA0 represent the continental levels. Different colors represent the different sub-regions. Note the differences in scales.

Fig.6. Soccer diagrams for the seasonal and geographical model performances in NA: a) winter, b) spring, c) summer and d) autumn. Mn and Md represent the mean and median ensembles, respectively. EU0 and NA0 represent the continental levels. Different colors represent the different sub-regions. Note the differences in scales.

Fig.7. Observed and simulated seasonal diurnal O_3 profiles in a) winter, b) spring, c) summer and d) autumn over EU2.

Fig.8. Observed and simulated seasonal diurnal O_3 profiles in a) winter, b) spring, c) summer and d) autumn over NA3.

Fig.9. Observed surface O_3 concentration bins against mean bias for the EU and NA domains for the O_3 season (May-September).

Fig.10. Observed and simulated (models, mean and median) vertical O_3 profiles averaged over 2010 in the EU domain. Note the differences in scales.

Fig.11. Observed and simulated (models, mean and median) vertical O_3 profiles averaged over 2010 in the NA domain. Note the differences in scales.





Fig.2.

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Fig.3.



Fig.4.



Fig.5.



Fig.6.



Fig.7.



Fig.8.









Fig.11.